

## Assessing Long Term Changes of Productivity in Some Floodplain Soils, Egypt, Using Spatial Analyses Techniques

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**S**OIL productivity evaluation for a long term has been a major hotspot in soil science. The current study aims at monitoring soil productivity within the last four decades in the middle of the Nile Delta, Egypt. To fulfill this objective, eleven soil profiles were collected to represent the different mapping units. Landsat ETM+ image dated at 2013 and digital elevation model (SRTM) were processed using ENVI 4.7 software to identify the main physiographic units in the studied area. The results indicate that flood plain is the main landscape, while the main landforms of the area under consideration are grouped as basins, mantles and river terraces. Two soil orders could be identified; Entisols and Aridisols that include four great groups, *i.e.*, *Typic Torrifuvents*, *Vertic Torrifuvent*, *Vertic Natrargids* and *Typic Natrargids*. The spatial analyst function in ArcGIS 9.3 was used to estimate the rating of moisture content, drainage condition, effective soil depth, texture/structure, soluble salt concentration, organic matter content, mineral exchange capacity/nature of clay and mineral reserve. According to soil productivity in 2013, results showed that 73.46% of the total area have decreased productivity compared with soil productivity in 1968 and this is due to many problems, and the most influencing factors are high soil salinity and high exchangeable sodium percentage and rising soil water table, while about 13.06% of these lands have increased their productivity and thus to improve the properties of these soils and increase the soil productivity, agricultural management must take care to reduce salinity, alkalinity, and improve the drainage system.

**Keywords:** Soil productivity, Physiographic units, Floodplain, Spatial analyses.

Increasing human population, farm fragmentation, and mismanagement of some farm services in developing countries decrease agricultural production. Soil productivity is the capacity of a soil in its natural environment to produce a specific plant or sequence of plants under specific systems of management inputs. For a long time, soil productivity evaluation has been a major hotspot in soil science (Agber and Ali, 2012). Productivity capacity is dependent on basic chemical and physical properties such as texture, pH and available water-holding capacity. This, coupled with differences in climate from one site to another, makes it difficult to evaluate differences in productivity capacity between soils or land parcels. Land productivity capacity or land quality is a comprehension, at

the same time a precise concept in terms of agricultural activities. It is defined as a measure of capability of land to perform specific functions (Devi and Kumar, 2008). The land productive capacity can be evaluated directly or indirectly. Direct evaluations are carried out in the field, greenhouses or laboratory by means of some experiments under given climatic and management conditions. Indirect evaluations consist basically in developing and applying models of varying complexity, thereby attempting to estimate land productivity (Delgado and Lopez, 1998 and Dengiz *et al.*, 2010). Riquier *et al.* (1970) suggested five rating for evaluating soil productivity, *i.e.*, extremely poor to nil V (0–7), poor IV (8–19), average III (20–34), good II (35–64) and excellent I (65–100).

The different types and procedures in land evaluation are gradually being developed. Interpreting soil qualities and site information for the agricultural use and management practices is integrated using geographical information system (FAO, 1991 and 2007). Eswaran *et al.* (2003) made use of Geographic Information System to assess and monitor quality of land. They combined the soil and climate variables as these influence agricultural productivity. GIS and RS offer a great potential to capture data through a variety of observation platforms and integrate them through their common spatial network. This advanced approach justifies the involvement of object-oriented database structures in the decision-making process as this digital framework is an efficient system for managing data records for easy access toward decision making (Adrian *et al.*, 2010). The spatial analysis was used in this study, it can be defined as the analytical techniques associated with the study of locations of geographic phenomena together with their spatial dimensions and their associated attributes (ESRI, 2009). Spatial analysis is useful for evaluating suitability, for estimating and predicting, and for interpreting and understanding the location and distribution of geographic features and phenomena. The use of spatial analyses techniques in evaluating the land capability allows producing multi-thematic maps and outlining the limiting factors, accordingly suitable suggestions could be attained to understand how to deal with these soils for sustainable agricultural use. Therefore, the aims of this study are to produce the physiographic map of the study area scale 1:50.000 reduced to the attached map, as well as to use spatial analyses techniques in monitoring the soil productivity in the flood plain soils of the middle Nile Delta in Egypt.

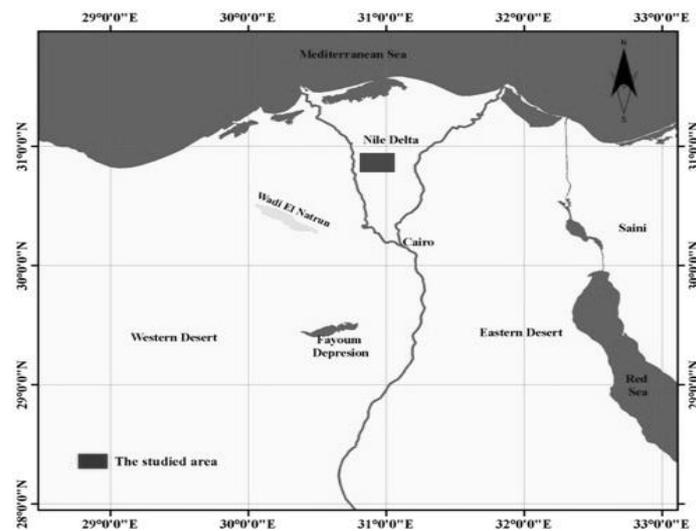
## Material and Methods

### *Location of the study area*

The study area occupies the middle part of Nile Delta, Egypt. It is bounded by 30°49'30" and 31°05'14" longitudes and 30°45'30" and 30°55'30" latitudes (Fig.1) covering a total area of 370.01 km<sup>2</sup> and has a population growth rate of nearly 2%. Based on Climatological Normal for Egypt (2011) and American Soil Taxonomy (USDA, 2010), the soil temperature regime of the studied area was defined as Thermic and soil moisture regime as Torric. The mean annual temperature reaches its maximum in June, July and August interval and dose

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exceeds 26.4°C but the temperature average reaches its minimum in January, February and March recording 13.7°C in Tanta station. The amount of annual rainfall is very low and mostly falls in winter; reaching about 3.8 mm/year. The studied area belongs to the late Pleistocene which is represented by the deposits of the neogene and those deposits accumulated during the recessional phases of the river (Said, 1993). Most of the study area is flat to almost flat (slope < 2.0%). The study area has been under intensive agricultural activities, whereas, rice, wheat, maize, cucumber and tomato with flood irrigation have been produced in the study area.



**Fig.1. Location of the study area.**

#### *Digital image processing and physiographic mapping*

Digital image processing was executed for Landsat ETM+ 8 satellite image (path 176, row 39) with spatial resolutions of 30 meters acquired in April 2013 using ENVI 4.7 software (ITT 2009). Data were calibrated to radiance using the inputs of image type, acquisition date and time. Images were stretched using linear 2%, smoothly filtered and their histograms were matched according to Lillesand and Kiefer (2007). Image was atmospherically corrected using FLAASH module (ITT 2009).

The landforms of the study area were initially determined from the satellite image and the digital elevation model (DEM) extracted from Shuttle Radar Topography Mission (SRTM) within Arc-GIS 9.3 software (ESRI, 2009). The extracted data generates a preliminary geomorphologic map which was checked and completed through 60 field observation points. Coverage was rectified and projected with projection (UTM), Zone (36) and Datum (WGS 84).

*Field studies and laboratory analyses*

Eleven soil profiles were collected and were taken from two sample areas, typically covering about 10% of the investigated area. This sample areas crossed the different mapping unites (Hengl and Rossiter, 2003). The locations of the soil profiles were selected to be the same sites previously studied by the Research Institute of Soils and Water (RISW 1968). Representative disturbed 44 soil samples have been collected. The samples were air dried and the less than 2 mm particles were used for chemical analyses. Electrical conductivity (EC) was determined conduct-metrically in saturated soil paste extracted. Cation exchange capacity (CEC) and exchangeable sodium percentage (ESP) were determined by ammonium acetate (NH<sub>4</sub>OAC). Organic matter (OM) was determined by the modified Walkley and Black method. Total calcium carbonate was determined volumetrically using Collin's calcimeter method. Particle size distribution of the soil samples was determined according to the international pipette method. The laboratory analyses were carried out using the soil survey laboratory methods manual (USDA, 2004). The soils were classified to the sub great groups level based on the American Soil Taxonomy (USDA, 2010). Then the physiographic and taxonomic units were correlated in order to identify the major soil sets of the studied area (Elberson and Catalan, 1987).

*Soil productivity assessment*

The land productivity classes were defined to monitoring soil productivity in the studied area using the rating and procedure after Riquier *et al.* (1970) model modified by FAO (2007). The system suggests the calculation of a productivity index considering eight factors as determining land productivity; Moisture (H), Drainage (D), Soil depth (P), Texture/Structure (T), Soluble salt concentration (S), Organic matter (O), Mineral exchange capacity (A) and Mineral reserve (M). These characteristics concern are rated and used to calculate the productivity index (PI) according to the following equation:

$$PI = (H/100 * D/100 * P/100 * T/100 * S/100 * O/100 * A/100 * M/100) * 100$$

Each factor is rated on a scale from 0 to 100 and the resultant index of productivity, also lying between 0 and 100, is set against a scale placing the soil in one or other of five productivity classes. Each of land and soil and land characteristics with associated attribute data are digitally encoded in a GIS database to eventually generate eight thematic layers. Based on comparing between the data extracted from RISW report, (1968) and the data resulting from this study the previous and actual soil productivity maps were produced.

## **Results and Discussion**

*Physiography and soils*

Field survey data, Landsat ETM images, and DEM were used to define the physiographic units in the study area as shown in Fig. 2 and Table 1. The obtained results indicate that the major landscape in the studied area is the flood plain. These soils are originated from Nile sediments before High Dam  
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construction. They are developed from sediments of Ethiopian plateau transported by Nile River and subsequently deposited in both the valley and Delta. The flood plain in the study area has been divided into four landforms are namely decantation basins, overflow basins, river terraces and overflow mantles. The correlation between physiography and soils was carried out, the produced data reveal that the soils of the main physiographic units in the area could be arranged under the landform level in the following:

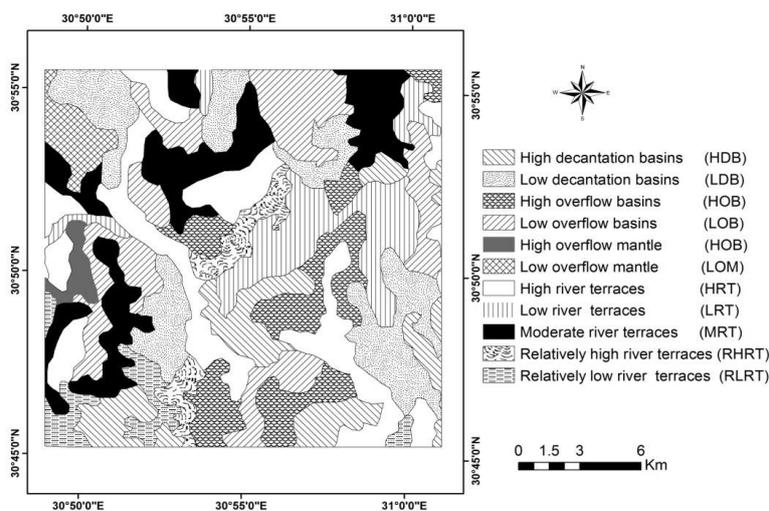


Fig.2. Physiography of the study area.

TABLE 1. Areas of the different physiographic units in the studied area.

Landscape	Relief/ Molding	Lithology/ Origin	Landform	Map unit	Area (Km <sup>2</sup> )	Area (%)
Flood plain	Flat to almost flat	Alluvial deposits	<b>Basins</b>			
			High decantation basins	HDB	35.59	9.62
			Low decantation basins	LDB	55.36	14.96
			High overflow basins	HOB	38.59	10.43
			Low overflow basins	LOB	53.41	14.43
			<b>Mantle</b>			
			High overflow mantle	HOM	4.65	1.26
			Low overflow mantle	LOM	9.33	2.52
			<b>River terraces</b>			
			High river terraces	HRT	59.67	16.13
			Relatively high river terraces	RHRT	12.83	3.47
			Moderate river terraces	MRT	49.88	13.48
			Relatively low river terraces	RLRT	12.73	3.44
			Low river terraces	LRT	37.97	10.26
				370.01	100	

*Soils of decantation basins*

Decantation basins are a low-lying area which is wholly or largely surrounded by higher land. It receives its ground deposits from overflow basins as it is lower than it. This landform covers an area of 86.4 km<sup>2</sup>, including soils of high decantation basins and soils of low decantation basin. These are classified to the sub-great group level as Typic Natrargids and Typic Torrfluvents, respectively. The soils of these units have different patterns of texture classes, where the soils of this unit vary from loam, clay loam and silty clay loam in the different profiles layers. The calcium carbonate content ranges between 2.09 and 4.16%. The high values are due to the occurrence of shell fragments. The organic matter content ranges between 0.49 and 3.83% in the successive layers of the studied soil profiles. The high values of organic matter in the surface and subsurface layers are related to the continuous adding of the organic manure. The electrical conductivity (ECe) values range between 4.39 and 21.65 dS/m in the successive layers of soil profiles. The cation exchange capacity (CEC) value is a result of the clay and organic matter content. However, the clay mineral plays a significant role, the cation exchange capacity ranges between 42.73 and 52.67 cmol+/kg soils. The exchangeable sodium percentage (ESP) values fluctuate between 19.36 and 37.68% in the different layers of the representative soil profiles.

*Soils of overflow basins*

Overflow basins are large, bowl-shaped depression in the surface of the land floor. It deposits its load of sedimentation carried by water into decantation basins. The total area of this landform is 83.64 km<sup>2</sup>, including high overflow basins and low overflow basin. These are classified to the sub-great group level as Vertic Torrfluvents and Typic Natrargids, respectively. The texture class is clay loam and loam in the different profile layers. The calcium carbonate content ranges between 1.5 and 4.2 %. The organic matter content ranges between 0.63 and 2.93% in the successive layers of the studied soil profiles. The EC values range between 4.62 and 22.45dS/m in the successive layers of soil profiles. The cation exchange capacity ranges between 42.73 and 46.32 cmol+/kg soils. The exchangeable sodium percentage (ESP) values fluctuate between 22.34 and 29.81% in the different layers of the representative soil profiles.

*Soils of river terraces*

These soils represent the old deltaic plain of recent age and occur at the edge of the decantation basins. These soils are formed of terraces at various heights above the valley floor. These land forms cover an area of 185.99 km<sup>2</sup>, including high, relatively high, moderate, relatively low and low river terraces. These are classified to the sub-great group level as Vertic natriargids and Typic Torrfluvents. These soils have different patterns of texture classes, where the soils of this unit vary from loam to silt clay loam in the different profile layers. The calcium carbonate content ranges between 0.53 and 4.28 %. The organic matter content ranges between 0.47 and 2.87% in the successive layers of the studied soil profiles. The EC values range between 2.94 and 11.69 dS/m. The cation exchange capacity ranges between 27.74 and 47.92 cmol+/kg soils. The exchangeable sodium percentage (ESP) values fluctuate between 6.87 and 27.15% in the different layers of the representative soil profiles.

*Soils of overflow mantles*

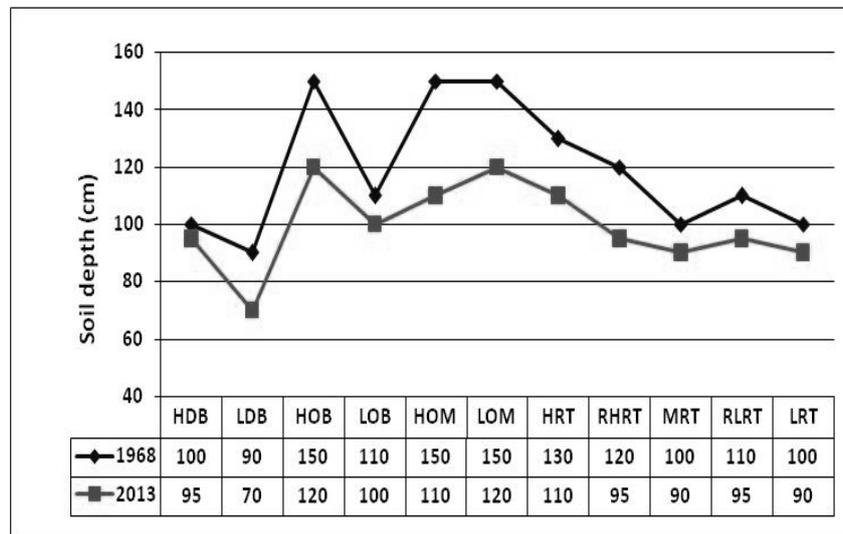
Soils of overflow mantles cover an area of 13.98 km<sup>2</sup>, including high and low overflow mantles. These are classified to the sub-great group level as Vertic Natrargids and Typic Natrargids. The soils of this unit vary from silt loam and clay in the different profile layers. The calcium carbonate content is low ranging between 1.28 and 2.88 %. The organic matter content records a range of 0.67 and 1.34%. The EC values are moderate to high ranging between 8.71 and 18.40 dS/m. Cation exchange capacity is high where it ranges between 40.69 and 49.35 cmol+/kg soils. Exchangeable sodium percentage is very high where it ranges between 23.88 and 29.98%.

*Monitoring of soil characteristics in the studied area*

Eleven soil profiles were investigated during 2013, these profiles were collected from the same locations previously studied in 1968. The changes in soil characteristics based on some selected soil properties such as soil depth, electrical conductivity (EC), exchangeable sodium percent (ESP), organic matter (OM), cation exchange capacity (CEC) and texture as shown in Table 2 and Fig. 3-7. As shown in Fig. 3, the water table level in all studied mapping units increased compared with the water table level in 1968. The decrease of soil depth in the study area may be due to the imbalance between irrigation and drainage, also, the use of flood irrigation and intensive agriculture may increase the water table level (El Baroudy, 2011). Figure 4 shows the change in organic matter content from 0.68–1.76% in 1968 to 0.93–3.65% in 2013 due to the addition of organic manure to soil and continues planting after 1964 may lead to increase the organic matter in the soil. Generally, cation exchange capacity (CEC) increased in the studied mapping units (Fig. 5), this increasing due to increase of organic matter content where the CEC value is a result of the clay and organic matter content (Gab-Allah, 2013). Figure 6 shows that the electrical conductivity (EC) in all mapping units increased from 3.57-8.32 dS/m (slightly salinity to moderately salinity) in 1968 to 5.94-16.32 dS/m (moderately salinity to very high salinity) in 2013. This increasing of EC values can be a result of poor management of irrigation schemes, whereas the use of flood irrigation and intensive agriculture may increase the water table level and hence increase the salinity of the soil surface (Ali, 2012 and Shalaby, 2013). Also, Fig. 7 shows that exchangeable sodium percent (ESP) increased in the studied area from 8.64-26.27 % (low alkalinity to high alkalinity) in 1968 to 12.63-28.41 % (moderately alkalinity to high alkalinity) in 2013. In this context El Baroudy (2010) reported that for the human induced salinization and alkalization in the middle Nile Delta can be result of the two causes, firstly, it can be the result of poor management of irrigation schemes. A high salt content of the irrigation water or too little attention given to the drainage of irrigated fields can be easily lead to rapid salinization. This type of salt accumulation mainly occurs under arid and semi-arid condition. A second type occurs where human activities lead to an increase in evapo-transpiration of soil moisture in areas of high salt-containing parent materials or with saline groundwater.

**TABLE 2. Monitoring of the main land characteristics in the studied area.**

Profile No.	Mapping unit	Water table level (cm)		EC dS/m		ESP %		OM %		CEC Cmol+/kg soil		Texture	
		1968	2013	1968	2013	1968	2013	1968	2013	1968	2013	1968	2013
1	HDB	100	95	4.32	5.94	8.64	12.63	1.65	3.65	47.68	49.25	Clay loam	
2	LDB	90	70	7.84	14.15	14.74	20.47	0.98	2.47	40.35	42.62	Silt clay loam	
3	HOB	150	120	7.4	14.5	14.83	20.47	0.68	2.31	42.35	44.25	Loam	
4	LOB	110	100	8.32	16.32	16.84	26.7	0.91	1.45	43.19	43.67	Loam	
5	HOM	150	110	6.73	12.44	16.24	27.43	0.96	1.36	56.89	58.36	Clay	
6	LOM	150	120	7.65	13.91	17.57	26.75	0.94	2.21	41.25	42.45	Silt loam	
7	HRT	130	110	3.57	9.11	24.26	26.65	1.83	2.16	43.26	44.68	Loam	
8	RHRT	120	95	4.28	6.11	26.27	28.41	1.76	2.14	37.15	39.71	Silt loam	
9	MRT	100	90	5.35	7.62	25.35	27.25	0.92	1.84	38.99	39.54	Silt clay	
10	RLRT	110	95	5.74	6.58	10.58	20.43	0.99	2.26	43.52	45.32	Clay loam	
11	LRT	100	90	7.42	9.35	12.65	20.43	0.89	0.93	28.74	29.54	Loam	

**Fig. 3. Changes of soil depth in the different landforms between 1968 and 2013.**

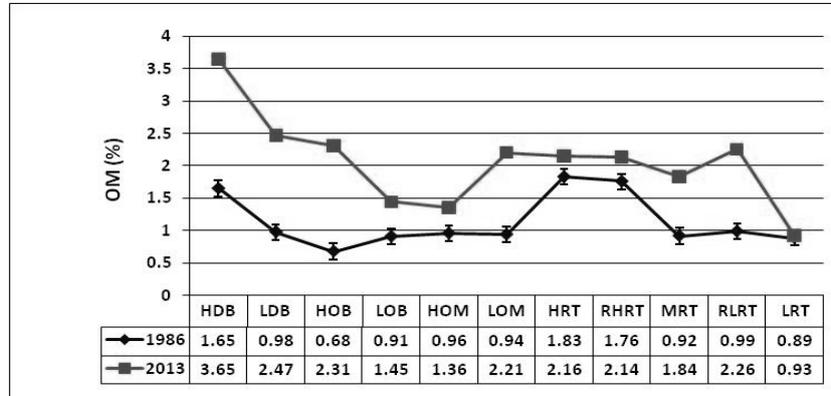


Fig.4. Changes of OM in the different landforms between 1968 and 2013.

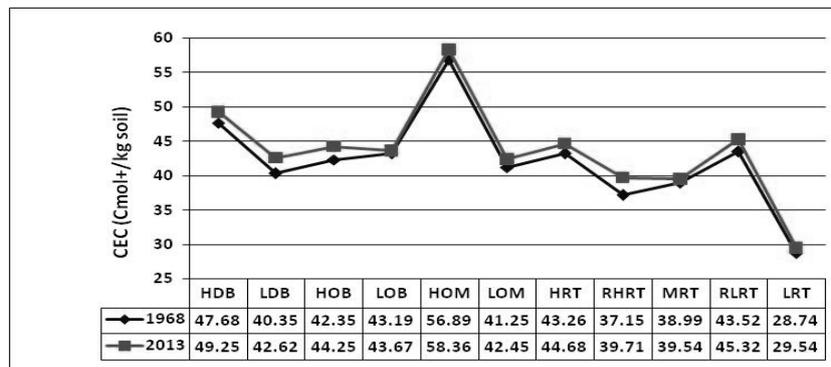


Fig.5. Changes of CEC in the different landforms between 1968 and 2013.

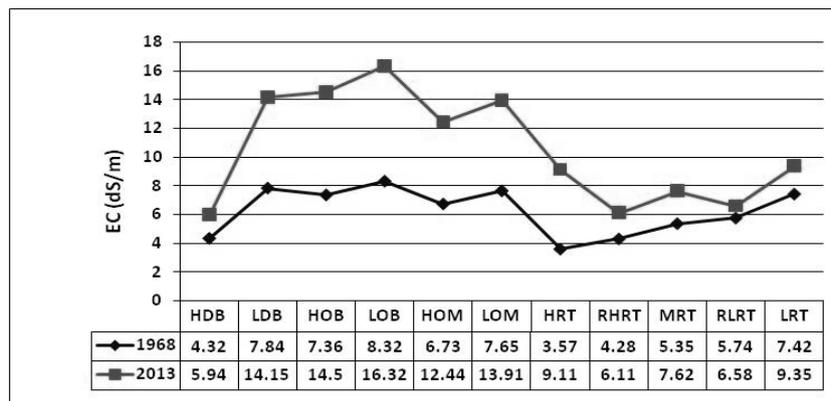


Fig.6. Changes of soil salinity in the different landforms between 1968 and 2013.

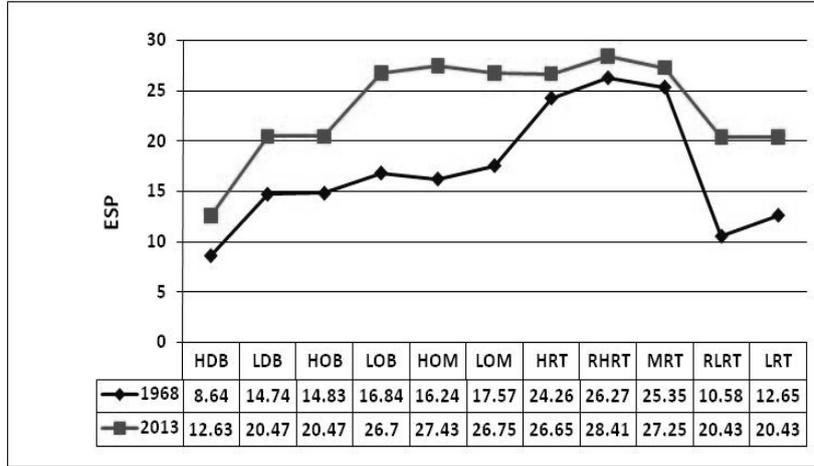


Fig.7. Changes of ESP in the different landforms between 1968 and 2013.

#### Monitoring of soil productivity in the studied area

A comparative study was carried out to monitor changes in soil productivity based on some selected soil properties defined by the productivity index as shown in Tables 3 and 4 and Fig. 8. Soil characteristics and productivity criteria are matched for getting soil productivity. Monitoring of soil productivity includes the following main landforms, *i.e.*, decantation basins, overflow basins, river terraces and overflow mantle. Table 4 represents the changes in area by square kilometers for each soil productivity grad (G) and index (PI) in the study area. The sites selection depends mainly upon previous study carried out by RISW (1968) to make comparison as well as monitoring more realistic. The soil characteristics of previous and current studies were grouped and recalculated to meet the requirements of Riquier *et al.* (1970) modified by FAO (2007) as follows:

#### Soil productivity changes in the decantation basins

This study demonstrated that soils of decantation basins are classified as good grades (II) in 1968. The productivity index (PI) of the high decantation basins (HDB) changed positively from good grade (55.40) to excellent grade (68.40), the improvement of soil productivity in this landform due to the good quality of land management practice (the continuously added organic manure, conservation and improvement drainages). The calculated values of PI for the low decantation basins (LDB) unit were shifted from good grade (41.9) in 1968 to poor grade (9.23) in 2013, the main factors reduced the productivity index in LDB are soluble salts concentration, compaction, alkalinity and the effective soil depth.

**TABLE 3. Soil productivity index in the studied area during 1968 and 2013.**

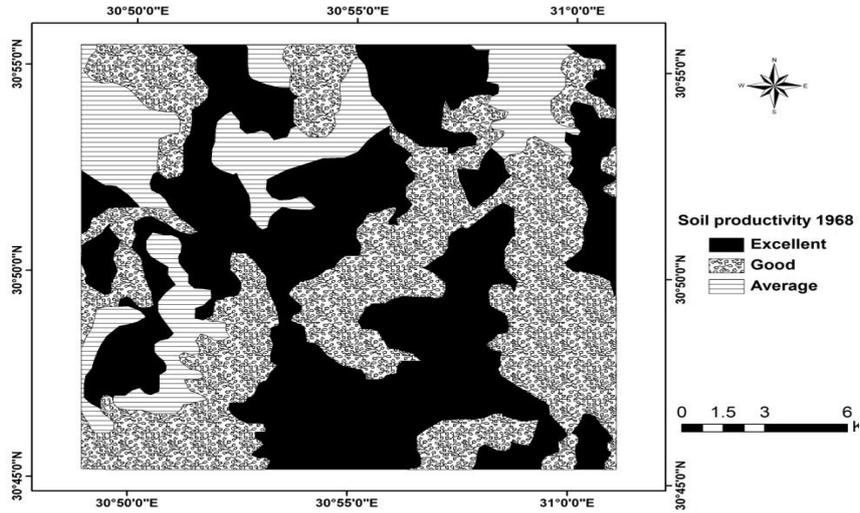
Mapping unit	H		D		P		T		S		O		A		M		Productivity index	
	1968	2013	1968	2013	1968	2013	1968	2013	1968	2013	1968	2013	1968	2013	1968	2013	1968	2013
HDB	100	100	80	100	90	100	90	90	100	80	90	100	100	100	95	95	55.40	68.40
LDB	100	100	80	90	90	80	90	90	80	15	85	100	100	100	95	95	41.86	9.23
HOB	100	100	100	90	100	100	100	100	80	15	85	100	100	100	95	95	64.60	12.83
LOB	100	100	90	90	90	100	100	100	100	15	85	90	100	100	100	100	68.85	12.15
HOM	100	100	100	90	100	100	60	60	80	25	85	90	100	100	95	95	38.76	11.54
LOM	100	100	80	100	100	100	100	100	40	15	85	100	100	100	100	100	27.20	15.00
HRT	100	100	80	90	100	100	100	100	100	40	90	100	100	100	100	100	72.00	36.00
RHRT	100	100	90	80	90	100	100	100	100	80	90	100	95	95	100	100	69.26	60.80
MRT	100	100	80	80	80	80	60	60	90	80	85	90	100	100	100	100	29.38	27.65
RLRT	100	100	90	90	90	100	90	90	90	80	85	100	100	100	100	100	55.77	64.80
LRT	100	100	80	80	90	80	100	100	80	40	85	85	100	100	100	100	48.96	21.76

Moisture (H), Drainage (D), Soil depth (P), Texture/Structure (T), Soluble salt concentration (S), Organic matter (O), Mineral exchange capacity (A) and Mineral reserve (M).

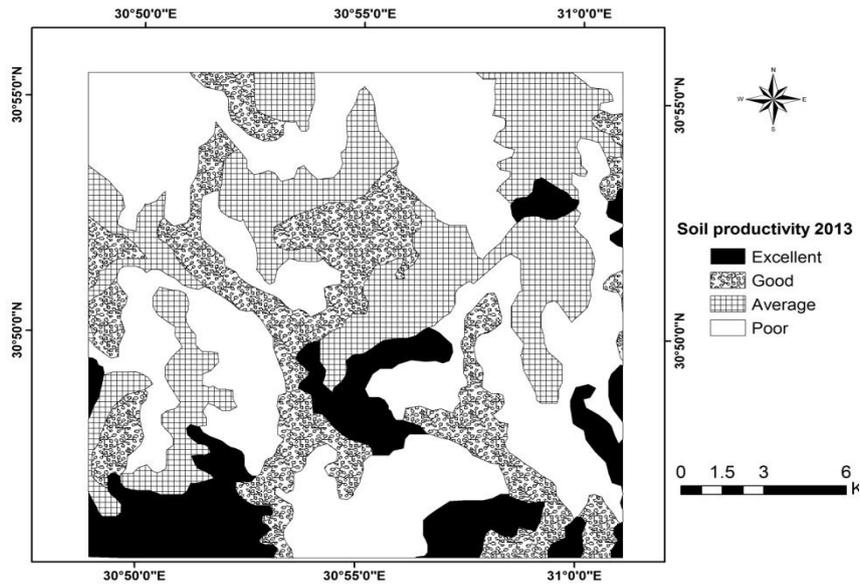
**TABLE 4. Areas of soil productivity grades in 1968 and 2013.**

PI	Grade	1968		2013	
		Area (km <sup>2</sup> )	%	Area (km <sup>2</sup> )	%
0-7	Nil to extremely poor (V)	-	-	-	-
8-19	Poor (IV)	-	-	161.34	43.61
20-34	Average (III)	59.21	16.00	87.85	23.74
35-64	Good (II)	146.30	39.54	72.50	19.59
65-100	Excellent (I)	164.50	44.46	48.32	13.06

(a)



(b)



**Fig. 8. Productivity index of the studied area, (a) 1968 and (b) 2013.**

*Soil productivity changes in the overflow basins*

According to the results, this landform includes high overflow basins (HOB) and low overflow basins (LOB). Values of soil productivity index in the HOB and LOB landforms were strongly decreased from excellent grade (I) in 1968 to

poor grade (IV) in 2013. The reduction of soil productivity in these units is related to increase soil salinity (from moderate salinity to very high salinity), increase water table level and increase soil alkalinity (from moderate alkalinity to high alkalinity) and degraded drainage condition

#### *Soil productivity changes in the river terraces*

These soils are formed of terraces at various heights above the valley floor. These land forms including high river terraces (HRT), relatively high river terraces (RHRT), moderate river terraces (MRT), relatively low river terraces (RLRT) and low river terraces (LRT). Values of soil productivity index in the HRT and RHRT landforms were decreased from excellent grade (I) in 1968 to good grade (II) in 2013, where PI in the HRT and RHRT decreased from 72.0 and 69.26 to 36.0 and 60.8, respectively during 1968 and 2013 period. Soil productivity index in the LRT decreased from good grade (48.96) to average grade (21.76) during 1968 and 2013 period. Moderate river terraces were classified as average grade (III) in the two studied periods. The productivity index (PI) of the relatively low river terraces (RLRT) changed positively from good grade (55.77) to excellent grade (64.8), the improvement of soil productivity in this landform due to the good quality of land management practice.

#### *Soil productivity changes in the overflow mantles*

Soils of overflow mantles divided to high overflow mantles (HOM) and low overflow mantles (LOM). The productivity index (PI) of the high overflow basins (HOB) changed negatively from good grade (38.76) to poor grade (11.54) during the studied period. The calculated values of PI for the low overflow basins (LOB) unit were shifted from average grade (27.2) in 1968 to poor grade (15.0) in 2013. The main factors reduced the productivity index in HOB & LOB are soluble salts concentration and the effective soil depth, where the soil salinity increased from 6.73-12.44 and 7.65 - 13.91 dS/m in HOB and LOB, respectively.

### **Conclusion**

Results from this study showed that remote sensing data and GIS could provide an effective approach to produce the physiographic–soil map and monitoring the soil productivity with reference to their spatial extent and distribution. The obtained thematic layers in the database will be of great help and basic sources for the planners and decision makers in sustainable planning and monitoring the changes in soil productivity. The study indicates that the productivity index was decreased in 73.46 % and was increased in 13.06 % of the study area between 1968 and 2013. The soil productivity in vast areas of flood plain was reduced due to the poor land management. Consequently, the agriculture development in the study area requires improving the drainage networks and proper land management that can be performed by governmental support and farmers themselves.

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## تقييم تغيرات إنتاجية بعض الاراضى الفيضية فى مصر على المدى الطويل باستخدام تقنيات التحليل المكاني

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يعتبر تقييم إنتاجية التربة على المدى الطويل من النقاط الهامة والرئيسية في المجال الزراعى. وتهدف الدراسة الحالية إلى رصد إنتاجية التربة خلال العقود الأربعة الماضية في وسط دلتا النيل، مصر. ولتحقيق هذا الهدف تم اخذ 11 قطاعا أرضيا من منطقة الدراسة لتمثيل الوحدات الارضية . واستخدمت صورة القمر الصناعى Landsat ETM+ لعام 2013 وتم معالجة بيانات الماسح الرادارى الطبوغرافى (SRTM) للحصول على النموذج ثلاثى الابعاد (DEM) باستخدام برنامج ENVI وذلك لإنتاج الخريطة الفسيوجرافية لمنطقة الدراسة.

وأوضحت النتائج المتحصل عليها أن الأحواض والمائل والشرفات النهرية هي الوحدات التضاريسية الرئيسية الموجودة فى الاراضى الفيضية المدروسة وتم تقسيمها الى رتبتين اساسيتين هما رتبة Entisols ورتبة Aridisols والتي يمثلها اربع مجموعات عظمى هي Vertic ،Typic Torrfluvents و Vertic Natrargids ،Torrifluent و Typic Natrargids .

وقد استخدمت التحليلات المكانية كأحد ادوات نظم المعلومات الجغرافية من خلال برنامج Arc-GIS 9.3 لحساب معاملات كلا من المحتوى الرطوبى ، وحالة الصرف ، وعمق الماء الأرضى ، و القوام /البناء ، وتركيز الاملاح الذائبة ، ومحتوى الأرض من المادة العضوية والسعة التبادلية المعدنية وذلك لتقييم إنتاجية التربة فى منطقة الدراسة.

وعند مقارنة نتائج إنتاجية التربة عام 2013 مع إنتاجية التربة عام 1968 أوضحت النتائج أن 73.46% من مساحة الاراضى الكلية المدروسة قد انخفضت إنتاجيتها وهذا راجع الى مشاكل عديدة ، واكثر هذه العوامل تأثيرا هي ارتفاع ملوحة التربة وارتفاع نسبة الصوديوم المتبادل وارتفاع منسوب الماء الارضى ، بينما حوالى 13.06% من هذه الاراضى قد زادت إنتاجيتها وبالتالي لتحسين خواص هذه الاراضى وزيادة إنتاجيتها يجب العناية بالادارة الزراعية لتقليل الملوحة ، القلوية ، وتحسين نظام الصرف.